The hydrostatic approximation for stratified fluids An open problem

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Outline



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e problem with hydrostatic equations

A result with eddy diffusivity

Benefits from isopycnal coordinates

The initial set of equations

Motivated by geophysical flows, we study the heterogeneous, incompressible Euler equations with gravity force:¹

$$\begin{aligned} \partial_t \rho + \nabla_{\mathbf{x}} \cdot (\rho \mathbf{u}) + \partial_z (\rho w) &= 0, \\ \rho (\partial_t \mathbf{u} + (\mathbf{u} \cdot \nabla_{\mathbf{x}}) \mathbf{u} + w \partial_z \mathbf{u}) + \nabla_{\mathbf{x}} P &= 0, \\ \rho (\partial_t w + \mathbf{u} \cdot \nabla_{\mathbf{x}} w + w \partial_z w) + \partial_z P + g \rho &= 0, \\ \nabla_{\mathbf{x}} \cdot \mathbf{u} + \partial_z w &= 0, \\ \text{boundary conditions.} \end{aligned}$$

The velocity field is denoted $(\mathbf{u}, w) : \Omega \to \mathbb{R}^d \times \mathbb{R}$, the density $\rho : \Omega \to \mathbb{R}^+_*$. The pressure $P : \Omega \to \mathbb{R}$ is reconstructed through the elliptic equations

 $-\nabla_{\mathbf{x}} \cdot \left(\frac{1}{\rho} \nabla_{\mathbf{x}} P\right) - \partial_{z} \left(\frac{1}{\rho} \partial_{z} P\right) = \nabla_{\mathbf{x}} \cdot \left((\mathbf{u} \cdot \nabla_{\mathbf{x}}) \mathbf{u} + w \partial_{z} \mathbf{u} \right) + \partial_{z} \left(g \rho + \mathbf{u} \cdot \nabla_{\mathbf{x}} w + w \partial_{z} w \right).$

Using the hydrostatic approximation, the system becomes ...

¹We are discarding the Coriolis force because we are interested in short-time stability. We are also discarding (for now) eddy viscosity and diffusivity.

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The hydrostatic approximation

The heterogeneous, incompressible Euler equations with gravity force

$$\partial_{t}\rho + \nabla_{\mathbf{x}} \cdot (\rho \mathbf{u}) + \partial_{z}(\rho w) = 0,$$

$$\rho(\partial_{t}\mathbf{u} + (\mathbf{u} \cdot \nabla_{\mathbf{x}})\mathbf{u} + w\partial_{z}\mathbf{u}) + \nabla_{\mathbf{x}}P = 0,$$

$$\rho(\partial_{t}w + \mathbf{u} \cdot \nabla_{\mathbf{x}}w + w\partial_{z}w) + \partial_{z}P + \rho = 0,$$

$$\nabla_{\mathbf{x}} \cdot \mathbf{u} + \partial_{z}w = 0,$$

boundary conditions.
(E)

with the hydrostatic approximation becomes

$$\begin{cases} \partial_t \rho + \nabla_{\mathbf{x}} \cdot (\rho \mathbf{u}) + \partial_z (\rho w) = 0, \\ \rho (\partial_t \mathbf{u} + (\mathbf{u} \cdot \nabla_{\mathbf{x}}) \mathbf{u} + w \partial_z \mathbf{u}) + \nabla_{\mathbf{x}} P = 0, \\ \partial_z P = -\rho, \qquad \Longrightarrow \qquad P = P|_{z=z_{\text{surf}}} + \int_{\cdot}^{z_{\text{surf}}} \rho \, dz \quad (\mathsf{H}) \\ \partial_z w = -\nabla_{\mathbf{x}} \cdot \mathbf{u}, \qquad \Longrightarrow \qquad w = w|_{z=z_{\text{bot}}} - \int_{z_{\text{bot}}}^{\cdot} \nabla_{\mathbf{x}} \cdot \mathbf{u} \, dz \\ \text{boundary conditions.} \end{cases}$$

Qn: stability of shear flows $(\rho, \mathbf{u})(t, \mathbf{x}, z) = (\underline{\rho}(z), \underline{\mathbf{u}}(z))$?



The problem with hydrostatic equations

(partial and biased) state of the art on the hydrostatic equations

Homogeneous case: $\rho \equiv 1$.

- Spectral stability of the linearized system about shear flows, under the Rayleigh criterion $\underline{\mathbf{u}}''(z) \neq 0$. [Rayleigh (1880)][Arnold][Drazin&Reid]
- Ill-posedness of the linearized system about certain shear flows. [Renardy '09]
- Well-posedness of the (nonlinear) system in Sobolev spaces under the Rayleigh criterion. [Masmoudi&Wong '12] (after [Grenier '99], [Brenier '03])

Stably stratified case: $\partial_z \rho < 0$.

• Spectral stability of the linearized system about shear flows, under the Miles and Howard criterion $\frac{1}{4}|\underline{\mathbf{u}}'(z)|^2 \leq \frac{-\underline{\rho}'(z)}{\rho(z)}$. [Miles '61][Howard '61]

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Open problem

Well-posedness of the initial-value problem for the (linear or nonlinear) system?

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The problem with hydrostatic equations $\bigcirc \bigcirc \bigcirc$

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The linearized system about shear flows

Let us linearize the system

$$\begin{aligned} &\langle \partial_t \rho + \nabla_{\mathbf{x}} \cdot (\rho \mathbf{u}) + \partial_z (\rho w) = 0, \\ &\rho (\partial_t \mathbf{u} + (\mathbf{u} \cdot \nabla_{\mathbf{x}}) \mathbf{u} + w \partial_z \mathbf{u}) + \nabla_{\mathbf{x}} P = 0, \\ &\partial_z P = -\rho, \\ &\partial_z w = -\nabla_{\mathbf{x}} \cdot \mathbf{u}, \\ &\text{boundary conditions (periodic).} \end{aligned}$$
(H)

about shear solutions $(\rho, \mathbf{u})(t, \mathbf{x}, z) = (\underline{\rho}(z), \underline{\mathbf{u}}(z))$. We get

$$\begin{cases} \partial_t \rho + \underline{\mathbf{u}}(z) \cdot \nabla_{\mathbf{x}} \rho + \underline{\rho}'(z) w = 0, \\ \partial_t \mathbf{u} + (\underline{\mathbf{u}}(z) \cdot \nabla_{\mathbf{x}}) \mathbf{u} + \underline{\mathbf{u}}'(z) w + \frac{1}{\underline{\rho}(z)} \nabla_{\mathbf{x}} P = 0, \\ \partial_z P = -\rho, \qquad \qquad (\Longrightarrow P = \int_{\cdot}^1 \rho \, \mathrm{d}z \stackrel{\mathrm{def}}{=} \mathsf{L}\rho) \qquad (\mathsf{L}) \\ \partial_z w = -\nabla_{\mathbf{x}} \cdot \mathbf{u}, \qquad (\Longrightarrow w = -\int_0^{\cdot} \nabla_{\mathbf{x}} \cdot \mathbf{u} \, \mathrm{d}z = -\mathsf{L}^* \nabla_{\mathbf{x}} \cdot \mathbf{u}) \\ \text{boundary conditions (periodic).} \end{cases}$$

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about shear solutions $(\rho, \mathbf{u})(t, \mathbf{x}, z) = (\underline{\rho}(z), \underline{\mathbf{u}}(z))$. We get

$$\begin{cases} \partial_t \rho + \underline{\mathbf{u}}(z) \cdot \nabla_{\mathbf{x}} \rho - \underline{\rho}'(z) \mathsf{L}^* \nabla_{\mathbf{x}} \cdot \mathbf{u} = \mathbf{0}, \\ \partial_t \mathbf{u} + (\underline{\mathbf{u}}(z) \cdot \nabla_{\mathbf{x}}) \mathbf{u} - \underline{\mathbf{u}}'(z) \mathsf{L}^* \nabla_{\mathbf{x}} \cdot \mathbf{u} + \frac{1}{\underline{\rho}(z)} \nabla_{\mathbf{x}} \mathsf{L} \rho = \mathbf{0}, \\ \text{boundary conditions (periodic).} \end{cases}$$
(L)

The linearized system enjoys a symmetric structure if $\rho'(z) < 0$, $\mathbf{\underline{u}}'(z) = \mathbf{0}$.

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The linearized system about shear flows

$$\begin{aligned} \partial_t \rho + \underline{\mathbf{u}}(z) \cdot \nabla_{\mathbf{x}} \rho - \underline{\rho}'(z) \mathsf{L}^* \nabla_{\mathbf{x}} \cdot \mathbf{u} &= 0, \\ \partial_t \mathbf{u} + (\underline{\mathbf{u}}(z) \cdot \nabla_{\mathbf{x}}) \mathbf{u} - \underline{\mathbf{u}}'(z) \mathsf{L}^* \nabla_{\mathbf{x}} \cdot \mathbf{u} + \frac{1}{\underline{\rho}(z)} \nabla_{\mathbf{x}} \mathsf{L} \rho &= 0, \end{aligned}$$
(L) boundary conditions (periodic).

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The linearized system about shear flows

$$\begin{aligned} \partial_t \rho + \underline{\mathbf{u}}(z) \cdot \nabla_{\mathbf{x}} \rho - \underline{\rho}'(z) \mathsf{L}^* \nabla_{\mathbf{x}} \cdot \mathbf{u} &= 0, \\ \partial_t \mathbf{u} + (\underline{\mathbf{u}}(z) \cdot \nabla_{\mathbf{x}}) \mathbf{u} - \underline{\mathbf{u}}'(z) \mathsf{L}^* \nabla_{\mathbf{x}} \cdot \mathbf{u} + \frac{1}{\underline{\rho}(z)} \nabla_{\mathbf{x}} \mathsf{L} \rho &= 0, \end{aligned}$$
(L) boundary conditions (periodic).

The linearized system enjoys a symmetric structure if $\underline{\rho}'(z) < 0$, $\underline{\mathbf{u}}'(z) = \mathbf{0}$.



The symmetric structure is ruined if $\underline{\mathbf{u}}'(z) \neq \mathbf{0}$.

$$\frac{\mathrm{d}}{\mathrm{d}t} \left(\left(\rho , \frac{-1}{\underline{\rho}'(z)} \rho \right)_{L^2_{\mathbf{x},z}} + \left(\mathbf{u} , \underline{\rho}(z)\mathbf{u} \right)_{L^2_{\mathbf{x},z}} \right) = \left(\mathbf{u} , \underline{\rho}(z)\underline{\mathbf{u}}'(z)\mathbf{L}^* \nabla_{\mathbf{x}} \cdot \mathbf{u} \right)_{L^2_{\mathbf{x},z}}.$$

$$\rightsquigarrow \text{ no (obvious) control of an energy.}$$

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Conclusion

$$\begin{cases} \partial_t \rho + \underline{\mathbf{u}}(z) \cdot \nabla_{\mathbf{x}} \rho - \underline{\rho}'(z) \mathsf{L}^* \nabla_{\mathbf{x}} \cdot \mathbf{u} = 0, \\ \partial_t \mathbf{u} + (\underline{\mathbf{u}}(z) \cdot \nabla_{\mathbf{x}}) \mathbf{u} - \underline{\mathbf{u}}'(z) \mathsf{L}^* \nabla_{\mathbf{x}} \cdot \mathbf{u} + \frac{1}{\underline{\rho}(z)} \nabla_{\mathbf{x}} \mathsf{L} \rho = 0, \end{cases}$$
(L) boundary conditions (periodic).

$$\frac{\mathrm{d}}{\mathrm{d}t} \left(\left(\rho , \frac{-1}{\underline{\rho}'(z)} \rho \right)_{L^2_{\mathbf{x},z}} + \left(\mathbf{u} , \underline{\rho}(z) \mathbf{u} \right)_{L^2_{\mathbf{x},z}} \right) = \left(\mathbf{u} , \underline{\rho}(z) \underline{\mathbf{u}}'(z) \mathsf{L}^* \nabla_{\mathbf{x}} \cdot \mathbf{u} \right)_{L^2_{\mathbf{x},z}}.$$

Possible high-frequency instabilities in the hydrostatic framework in the presence of shear velocity, $\underline{\mathbf{u}}'(z) \neq 0$, even under the Miles & Howard criterion. The stable stratification, $\rho'(z) < 0$, helps (a bit).

Possible ways around:

- Work in the analytic framework. e.g. [Kukavica, Temam, Vicol & Ziane '11]
- Add (horizontal) viscosity. e.g. [Cao, Li & Titi '16]
- Relax the hydrostatic approximation [Desjardins, Lannes & Saut '21]

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The problem with hydrostatic equation

with

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Eddy viscosity and diffusivity

The system studied in [Cao, Li & Titi '16] (among other works) is

$$\begin{aligned} \partial_t \rho + \mathbf{u} \cdot \nabla_{\mathbf{x}} \cdot \rho + w \partial_z \rho &= \kappa \Delta_{\mathbf{x}} \rho, \\ \rho(\partial_t \mathbf{u} + (\mathbf{u} \cdot \nabla_{\mathbf{x}}) \mathbf{u} + w \partial_z \mathbf{u}) + \nabla_{\mathbf{x}} P &= \nu \Delta_{\mathbf{x}} \mathbf{u}, \\ \partial_z P &= -\rho, \quad \partial_z w = -\nabla_{\mathbf{x}} \cdot \mathbf{u}, \\ \text{boundary conditions (periodic).} \end{aligned}$$
 (H_{\nu,\keta)}

The <u>horizontal</u> eddy viscosity ($\nu > 0$) and diffusivity ($\kappa > 0$) approximate effective isopycnal viscosity and diffusivity:

[Gent & McWilliams '90] propose the following parameterization

$$\begin{cases} \partial_{t}\rho + (\mathbf{u} + \mathbf{u}_{\star}) \cdot \nabla_{\mathbf{x}}\rho + (w + w_{\star})\partial_{z}\rho = 0, \\ \rho(\partial_{t}\mathbf{u} + ((\mathbf{u} + \mathbf{u}_{\star}) \cdot \nabla_{\mathbf{x}})\mathbf{u} + (w + w_{\star})\partial_{z}\mathbf{u}) + \nabla_{\mathbf{x}}P = 0, \\ \partial_{z}P = -\rho, \quad \partial_{z}w = -\nabla_{\mathbf{x}} \cdot \mathbf{u}, \\ \text{boundary conditions.} \\ \mathbf{u}_{\star} = \kappa \partial_{z} \left(\frac{\nabla_{\mathbf{x}}\rho}{\partial_{z}\rho}\right) , \quad w_{\star} = -\kappa \nabla_{\mathbf{x}} \cdot \left(\frac{\nabla_{\mathbf{x}}\rho}{\partial_{z}\rho}\right). \end{cases}$$
(H_{\kappa})

The problem with hydrostatic equations

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Main result

$$\begin{cases} \partial_t \rho + (\mathbf{u} + \mathbf{u}_{\star}) \cdot \nabla_{\mathbf{x}} \rho + (w + w_{\star}) \partial_z \rho = 0, \\ \rho (\partial_t \mathbf{u} + ((\mathbf{u} + \mathbf{u}_{\star}) \cdot \nabla_{\mathbf{x}}) \mathbf{u} + (w + w_{\star}) \partial_z \mathbf{u}) + \nabla_{\mathbf{x}} P = 0, \\ \partial_z P = -\rho, \quad \partial_z w = -\nabla_{\mathbf{x}} \cdot \mathbf{u}, \\ \text{boundary conditions (free surface),} \\ \mathbf{u}_{\star} = \kappa \partial_z \left(\frac{\nabla_x \rho}{\partial_z \rho} \right), \quad w_{\star} = -\kappa \nabla_{\mathbf{x}} \cdot \left(\frac{\nabla_x \rho}{\partial_z \rho} \right). \end{cases}$$
(H_{\kappa})

[R. Bianchini & VD]

For sufficiently regular data satisfying the (strict) stable stratification assumption

 $-\partial_z \rho|_{t=0} \ge \alpha > 0$

and for any $\kappa > 0$, there exists a unique (classical) solution to (H_{κ}) on the time interval [0, T] with

$$T^{-1} = C \left(1 + \kappa^{-1} \left(\left| \underline{\mathbf{u}}' \right|_{L_{z}^{2}}^{2} + M_{0}^{2} \right) \right),$$

where M_0 is the size of the initial deviation from the shear flow equilibrium $(\rho(z), \underline{\mathbf{u}}(z))$, and C depends only on M_0 , α and the size of $(\rho(z), \underline{\mathbf{u}}(z))$.



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Main result

[R. Bianchini & VD]

For sufficiently regular data satisfying the (strict) stable stratification assumption

 $-\partial_z \rho|_{t=0} \geq \alpha > \mathbf{0}$

and for any $\kappa>0,$ there exists a unique (classical) solution to $({\sf H}_\kappa)$ on the time interval [0,T] with

$$T^{-1} = C \left(1 + \kappa^{-1} \left(\left| \underline{\mathbf{u}}' \right|_{L^{2}_{\tau}}^{2} + M_{0}^{2} \right) \right),$$

where M_0 is the size of the initial deviation from the shear flow equilibrium $(\rho(z), \underline{\mathbf{u}}(z))$, and C depends only on M_0 , α and the size of $(\rho(z), \underline{\mathbf{u}}(z))$.

Remarks.

- We do not use viscosity (only diffusivity), nor analytic data;
- The destabilizing role of the shear velocity is apparent;
- Stable stratification is used in a crucial way in our proof;

 \rightsquigarrow we make use of isopycnal coordinates.

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A result with eddy diffusivity

Benefits from isopycnal coordinates $\bigcirc \bullet \bigcirc$

Isopycnal coordinates

We define the variable $h(t, \mathbf{x}, r) > 0$ from the density $\rho(t, \mathbf{x}, z)$ through

 $h \stackrel{\text{def}}{=} -\partial_r \zeta, \qquad \rho(t, \mathbf{x}, \zeta(t, \mathbf{x}, r)) = r, \quad \zeta(t, \mathbf{x}, \rho(t, \mathbf{x}, z)) = z.$



The problem with hydrostatic equation

Benefits from isopycnal coordinates $\bigcirc \bullet \bigcirc$

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The system

$$\begin{cases} \partial_t \rho + (\mathbf{u} + \mathbf{u}_{\star}) \cdot \nabla_{\mathbf{x}} \rho + (w + w_{\star}) \partial_z \rho = 0, \\ \rho (\partial_t \mathbf{u} + ((\mathbf{u} + \mathbf{u}_{\star}) \cdot \nabla_{\mathbf{x}}) \mathbf{u} + (w + w_{\star}) \partial_z \mathbf{u}) + \nabla_{\mathbf{x}} P = 0, \\ \partial_z P = -\rho, \quad \partial_z w = -\nabla_{\mathbf{x}} \cdot \mathbf{u}, \\ \text{boundary conditions (free surface).} \end{cases}$$
(H_{\kappa})

reads in isopycnal coordinates

$$\begin{cases} \partial_t h + \nabla_{\mathbf{x}} \cdot (h\mathbf{u}) = \kappa \Delta_{\mathbf{x}} h, \\ r \Big(\partial_t \mathbf{u} + \big(\big(\mathbf{u} + \kappa \frac{-\nabla_{\mathbf{x}} h}{h} \big) \cdot \nabla_{\mathbf{x}} \big) \mathbf{u} \Big) + \nabla_{\mathbf{x}} \psi = 0, \end{cases}$$
(H_{\kappa})

where

$$\psi(t, \mathbf{x}, r) \stackrel{\text{def}}{=} \rho_0 \int_{\rho_0}^{\rho_1} h(t, \mathbf{x}, r') \, \mathrm{d}r' + \int_{\rho_0}^{r} \int_{r'}^{\rho_1} h(t, \mathbf{x}, r'') \, \mathrm{d}r'' \, \mathrm{d}r'.$$

he problem with hydrostatic equations

Benefits from isopycnal coordinates $\circ \circ \bullet$

The system in isopycnal coordinates

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(H_{\kappa})

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Remarks:

- The eddy diffusivity of [Gent & McWilliams '90] is nice and simple.
- The advection in the variable z (or r) has disappeared.
- The domain is flattened.
- The system is easily discretized (multilayer framework) [Adim, w.i.p.]
- The system enjoys a partial symmetric structure analogous as the one exhibited in the Eulerian coordinates, but the symmetry defect involves an extra derivative on ζ ^{def} = ∫_r h (and not u).

 \rightsquigarrow The proof of our main result is based on the energy method on (H_{κ}), using product, commutator, composition estimates in anisotropic Sobolev spaces.

Thoughts to go

- The well-posedness of the heterogeneous hydrostatic equations, without diffusivity or viscosity, is an open problem.
- Isopycnal coordinates are interesting for numerical <u>and also</u> theoretical analyses.

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Thank you for your attention